Topographic Control of Asynchronous Glacial Advances: A Case Study from Annapurna, Nepal

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Topographic control of asynchronous glacial advances: A case study from Annapurna, Nepal
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[1] Differences in the timing of glacial advances, which are commonly attributed to climatic changes, can be due to variations in valley topography. Cosmogenic 10Be dates from 24 glacial moraine boulders in 5 valleys define two age populations, late-glacial and early Holocene. Moraine ages correlate with paleoglacier valley hypsometries. Moraines in valleys with lower maximum altitudes date to the late-glacial, whereas those in valleys with higher maximum altitudes are early Holocene. Two valleys with similar equilibrium-line altitudes (ELAs), but contrasting ages, are <5 km apart and share the same aspect, such that spatial differences in climate can be excluded. A glacial mass-balance cellular automata model of these two neighboring valleys predicts that change from a cooler-drier to warmer-wetter climate (as at the Holocene onset) would lead to the glacier in the higher altitude catchment advancing, while the lower one retreats or disappears, even though the ELA only shifted by ~120 m. Citation: Pratt-Sitaula, B., D. W. Burbank, A. M. Heimsath, N. F. Humphrey, M. Oskin, and J. Putkonen (2011), Topographic control of asynchronous glacial advances: A case study from Annapurna, Nepal, Geophys. Res. Lett., 38, L24502, doi:10.1029/2011GL049940.

1. Introduction
[2] Glaciers are potentially valuable recorders of terrestrial climate change – delicately tuned to the combined effect of snow fall and temperature. The complex reality of glacial behavior, however, can hinder efforts to create reliable climatic reconstructions. Although regional climatic signals are commonly discernible, individual glaciers have been shown to advance and retreat asynchronously at local, regional, and global scales [e.g., Benn and Owen, 1998; Gillespie and Molinar, 1995]. Further understanding of the factors that lead to disparate glacial behavior will improve interpretation of past climate and prediction of future glacial change. Previous researchers have suggested that hypsometry plays a role in glacial asynchrony [e.g., Chenet et al., 2010; Kerr, 1993], but this study is the first to test the hypothesis rigorously with a combination of modeling and field work. Results from this study quantify, more thoroughly than in previous studies, how glacial hypsometry, as well as climate, can drive asynchronous glacial behavior.

2. Background and Methods
[3] Prior studies of glaciation in the Annapurna area are limited to broad regional surveys [e.g., Fort and Derbyshire, 1988; Hagen, 1968; Owen et al., 1998], recent glacial erosion estimates [Heimsath and McGlynn, 2008], and limited, scattered 10Be dating [Zech et al., 2009]. Further east in Nepal, other glacial dating studies [Finkel et al., 2003; Gayer et al., 2006; Owen et al., 2009] have found ice extent maxima at times broadly consistent with those presented here. This study in the Annapurna region of central Nepal (Figure 1) combines field observations, 10Be dating of glacial boulders, digital elevation model (DEM) analysis, and cellular automata modeling of glacial mass balance to test the hypothesis that identical climate change can cause coeval advance and retreat of neighboring glaciers in drainage basins that have significantly different hypsometries.

2.1. Cosmogenic 10Be Dating
[4] We dated moraines in valleys spanning a range of locations and altitudinal extents. Two valleys lie north (leeward) of the Annapurna Range (Figure 1b); one is on the range crest (Figure 1c); and two lie to the south (windward) (Figure 1d). Syaktan valley (Figure 1b) and Danfe valley (Figure 1d) descend from peaks ≥5500 m, whereas ridgelines in the other 3 valleys are <5200 m high. The dated glacial valleys are small (<8 km2) and not complicated by significant avalanching which can displace snow far below the climatic equilibrium [Harper and Humphrey, 2003]. Valley bottoms are free of obvious hummocky morphology and lack large headwalls, suggesting that the glacier surfaces were free of extensive debris cover. Each valley has only one major set of preserved moraines and no modern glaciers. The moraines have multiple crests in places, but soil development and morphology do not change significantly from one crest to the next, thereby suggesting deposition over a short interval. Of particular note, the two northern valleys (Figure 1b) lie <5 km apart and have the same aspect and climate. Four to seven moraine-crest boulders were sampled from each valley (Figure 1 and Table S1 in the auxiliary material).1 We prepared samples using standard procedures.
[e.g., Kohl and Nishiizumi, 1992] and measured the $^{10}$Be concentrations at Lawrence Livermore National Laboratory. Ages were calculated using v2.2 CRONUS-Earth online calculator [Balco et al., 2008]. Ages presented in this paper are from the Lal [1991] and Stone [2000] time-dependent scaling model. Results from other scaling models (Table S2) nearly all fall within the external error bars of the time-dependent Lal/Stone model (Figure 2a), thus differences in scaling models, although important to consider, do not affect the conclusions of the study.

2.2. Topographic Analysis: Glacial Hypsometry and Equilibrium Line Altitudes

Based on valley morphology and moraine location determined from air photo and field analysis, we digitized areas for each paleoglacier onto georeferenced topographic maps [Finnish Meteorological Institute, 2001], and glacier-bed hypsometry was calculated from a 3-arcsecond (~90-m) Shuttle Radar Topography Mission data (Figure 3). For each paleoglacier, the former ELA (equilibrium-line altitude where net snow mass input is balanced by the local net snow melt) was estimated using the calculated hypsometries and assuming an accumulation-area ratio of 0.60. Work in Nepal [Ageta and Higuchi, 1984; Benn and Lehmkuhl, 2000] and elsewhere [e.g., Benn et al., 2005; Porter, 2000] shows that the accumulation zone typically occupies 50–80% of a glacier’s area. Because use of other accumulation-area ratios [e.g., Kern and Laszlo, 2010] or other methods (e.g., area-altitude balance ratio [Benn et al., 2005; Rea, 2009]) changes the absolute ELA estimates somewhat, but does not alter our primary conclusions, we used a method and ratio in common use for inter-study comparability.

2.3. Glacial Mass Balance Model

A cellular automata model [Harper and Humphrey, 2003] was used to investigate how changes in glacial ELA and mass balance influence glacial dynamics. The model requires an input ELA, a mass-balance curve (change in mass gain or loss versus altitude), and topography of the entire glacial valley. From these, the net mass of snow input is calculated for each altitude. At each time-step, the ice is

![Figure 1. a) Shaded relief map of the field area with locations of detailed maps. Main Marsyandi valley contains numerous degraded, pre-Holocene glacial and landslide deposits more fully described in regional surveys [Fort and Derbyshire, 1988; Hagen, 1968; Owen et al., 1998]. b–d) Detailed maps of sample locations, $^{10}$Be exposure ages, moraine crests, and paleoglaciers extents.](image)

![Figure 2. Glacial chronologies and paleoclimate proxies. Vertical gray bands indicate the temporal extent of the Holocene intensified monsoon and the Bølling-Ållerød. a) Moraine boulder $^{10}$Be exposure ages from this study. Names and boxes refer to individual valleys (see Figure 1). The gray error bar shows the internal error from laboratory, AMS, and snow-shielding uncertainties. The thin black error bar shows the external error from the production rate scaling model. *Indicates single sample that was split for analysis and provides a measure of laboratory repeatability. b) Paleoprecipitation proxies: Arabian Sea total organic carbon [Schulz et al., 1998], Bengal Sea freshwater [Kudrass et al., 2001], Hulu [Wang et al., 2001] and Dongge [Yuan et al., 2004] cave records in gray, Tinta cave record [Sinha et al., 2005] in black. c) Paleotemperature proxies: South China Sea surface temperatures (SST) in black [Kienast et al., 2001] and gray [Oppo and Sun, 2005] and GISP2 ice core record [Alley et al., 1995].](image)
routed across the landscape according to two basic assumptions about ice behavior. If the surface slope is >35°, it is unstable and any snow avalanches to the next lowest cell. If the ice surface is <35°, the ice thickness varies in order to maintain a basal shear stress ($\tau_b$) of 1 bar [Nye, 1952; Paterson, 1994]. Thus,

$$\tau_b = \rho ghS \approx 1 \text{ bar},$$  

where $\rho =$ density, $g =$ force due to gravity, $h =$ ice thickness, and $S =$ ice surface slope. Any ice thickness causing stresses exceeding 1 bar will flow to the lowest adjacent cell. To match the values determined for modern Annapurna area glaciers [Harper and Humphrey, 2003], the upper boundary for significant snowfall was set to 6200 m, the mass balance was set to be constant through the accumulation area until just above the ELA, where it decreases at a constant rate of 8 m/yr per km of altitude through the ablation area. The bulk of the studied glacial valleys are below 6000 m, such that model results are insensitive to the upper snow boundary. Likewise, changing the mass balance gradient in the ablation zone by $\pm$6 m/yr/km does not significantly change the outcome. Although the model does not mimic ice flow exactly, it succeeds in predicting general ice extent in modern Himalayan glaciers [Harper and Humphrey, 2003].

3. Results and Interpretation

The $^{10}$Be exposure ages were consistent within each valley (Table 1 and Figure 2a), with no outliers. The dates fall into 2 distinct populations: 15–13 ka and ~8 ka. Kicho valley ages range from 13.3 to 15.1 ka with a sample mean of 14.1 ± 0.8 ka, Midim valley ages are 12.8–13.6 ka with a mean of 13.3 ± 0.3 ka, and Khudi valley ages are 13.2–14.6 ka with a mean of 14.1 ± 0.4 ka. The younger ages come from Syaktan valley with a range of 7.4–9.3 ka with a mean of 8.2 ± 0.8 ka and Danfe valley where ages are 7.5–8.3 ka with a mean of 8.0 ± 0.3 ka. We conclude that spatially
variable climate cannot account for the asynchrony, because glaciers to both the north (Figure 1b) and the southeast (Figures 1c and 1d) fall in both age populations. The northern glaciers are particularly notable, because they lie so close together (<5 km apart) and share the same aspect; yet, the last advance of each occurred at strikingly different times.

Table 1. $^{10}$Be Moraine Boulder Exposure Ages

<table>
<thead>
<tr>
<th>Sample Number</th>
<th>Age (ka)</th>
<th>Internal Error (ka)</th>
<th>External Error (ka)</th>
<th>Site Average and SD (ka)</th>
</tr>
</thead>
<tbody>
<tr>
<td>CRN-222</td>
<td>14.2</td>
<td>0.5</td>
<td>1.3</td>
<td>14.1 ± 0.8</td>
</tr>
<tr>
<td>CRN-223</td>
<td>13.3</td>
<td>0.4</td>
<td>1.2</td>
<td>Kicho</td>
</tr>
<tr>
<td>CRN-224</td>
<td>15.1</td>
<td>0.6</td>
<td>1.4</td>
<td>Midim</td>
</tr>
<tr>
<td>CRN-225</td>
<td>13.8</td>
<td>0.5</td>
<td>1.2</td>
<td>Syaktan</td>
</tr>
<tr>
<td>CRN-226</td>
<td>7.9</td>
<td>0.3</td>
<td>0.7</td>
<td>8.2 ± 0.8</td>
</tr>
<tr>
<td>CRN-227</td>
<td>9.3</td>
<td>0.3</td>
<td>0.8</td>
<td>Danfe</td>
</tr>
<tr>
<td>CRN-228</td>
<td>8.0</td>
<td>0.3</td>
<td>0.7</td>
<td></td>
</tr>
<tr>
<td>CRN-229</td>
<td>7.4</td>
<td>0.3</td>
<td>0.7</td>
<td></td>
</tr>
<tr>
<td>CRN-246</td>
<td>13.5</td>
<td>0.5</td>
<td>1.2</td>
<td>13.2 ± 0.3</td>
</tr>
<tr>
<td>CRN-247</td>
<td>13.6</td>
<td>0.5</td>
<td>1.3</td>
<td>Midim</td>
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<tr>
<td>CRN-248</td>
<td>12.8</td>
<td>0.6</td>
<td>1.3</td>
<td></td>
</tr>
<tr>
<td>CRN-249</td>
<td>12.9</td>
<td>0.7</td>
<td>1.2</td>
<td></td>
</tr>
<tr>
<td>CRN-250</td>
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<td>0.5</td>
<td>1.2</td>
<td></td>
</tr>
<tr>
<td>CRN-251</td>
<td>14.2</td>
<td>0.4</td>
<td>1.3</td>
<td>14.1 ± 0.4</td>
</tr>
<tr>
<td>CRN-252</td>
<td>14.6</td>
<td>0.4</td>
<td>1.3</td>
<td>Khudi</td>
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<tr>
<td>CRN-253a</td>
<td>13.9</td>
<td>0.5</td>
<td>1.3</td>
<td></td>
</tr>
<tr>
<td>CRN-253b</td>
<td>13.2</td>
<td>0.4</td>
<td>1.2</td>
<td></td>
</tr>
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<td>13.8</td>
<td>0.4</td>
<td>1.2</td>
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</tr>
<tr>
<td>CRN-343</td>
<td>7.9</td>
<td>0.3</td>
<td>0.7</td>
<td>8.0 ± 0.3</td>
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<tr>
<td>CRN-344</td>
<td>7.5</td>
<td>0.2</td>
<td>0.7</td>
<td>Danfe</td>
</tr>
<tr>
<td>CRN-345</td>
<td>8.0</td>
<td>0.3</td>
<td>0.7</td>
<td></td>
</tr>
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<td>CRN-347</td>
<td>8.3</td>
<td>0.3</td>
<td>0.7</td>
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<tr>
<td>CRN-348</td>
<td>8.4</td>
<td>0.3</td>
<td>0.7</td>
<td></td>
</tr>
<tr>
<td>CRN-349</td>
<td>7.9</td>
<td>0.3</td>
<td>0.7</td>
<td></td>
</tr>
<tr>
<td>CRN-440</td>
<td>8.3</td>
<td>0.3</td>
<td>0.7</td>
<td></td>
</tr>
</tbody>
</table>

*Ages were calculated using the CRONUS-Earth online calculator V2.2 with the time-integrated Lal/Stone production rate scaling model. Site maximums are in bold. Errors are ±1σ. Internal errors include carrier density, AMS, and snow shielding uncertainties. External errors reflect scaling model/nucleide production rate uncertainties. Corrections were made for atmospheric thickness, geomagnetic position, topographic shielding, sample thickness, and snow cover. CRN-253 was split for analysis; the two resulting ages were averaged before the site average was calculated. See Table S1 and auxiliary material for additional information.

Figure 4. Schematic diagram of net mass accumulation for two equal-sized glaciers with different bed hypsometries under two different climate scenarios. a) Net accumulations are sub-equal when ELA and mass input are lower (cooler, drier); b) When ELA and mass input are raised (warmer, wetter), net accumulation increases for the higher altitude glacier, but decreases for the lower glacier.
In particular, warmer but wetter conditions could cause glaciers with higher-altitude source areas to advance as lower-altitude glaciers recede. To test this hypothesis, we used the cellular automata model to explore relationships among ELA, mass balance in the accumulation area, and position of the glacial toe. Consistent with our observations, we assume no significant debris cover: a key factor given the potential for such cover to strongly influence glacial behavior [Scherler et al., 2010]. We simulated “late-glacial” conditions with an ELA of 4780 m and “early Holocene” conditions with an ELA of 4900 m. In the late-glacial simulations, glaciers fill both valleys and reach a steady-state position by 300 model years (Figure 3c). With an ELA of 4780 m and a mass balance (water equivalent) in the accumulation zone of 0.45 m/yr, the toe of the model Kicho glacier reached the same level (4585 m) as the actual glacier did at its maximum (Figure 3b). Under the same conditions, the model predicts that the Syaktan glacier toe remains some 100 m above its lowest position. However, when the mass balance in the accumulation zone is increased as the ELA rises (paleoclimate indicators suggest that the Early Holocene was wetter – Figure 2), Syaktan actually advances, while Kicho glacier melts away (Figure 3d). A mass balance of 1.15 m/yr in the accumulation zone brings the Syaktan toe to its mapped altitude (4105 m) (Figure 3b).

The model is probably not so well constrained that the specific accumulation-zone mass balances (1.15 m/yr or 0.45 m/yr) are accurate, but the sense of change is significant. Thus, we interpret the results from moraine dating and computer modeling to show that a change in climate – in this case from a colder, drier time when the ELA and accumulation-zone mass balance were lower to one that was warmer and wetter when the ELA and mass balance were (slightly) higher – can lead some glaciers to advance, while others retreat, due solely to differences in their hypsometries (Figure 4). The net outcome, as recorded by moraine ages that differ by >40%, is asynchronous maximum ice extent in neighboring valleys.

4. Conclusions

The unique combination of factors – multiple, well-dated paleoglaciers that have different hypsometries and fall into two distinct age populations from different climatic periods – allowed us to test the effect of hypsometry on glacial dynamics. The results quantify how topography can play a fundamental role in glacial behavior, causing some glaciers to advance while others retreat in the face of uniform climate change. The study is the first of its kind and has several significant implications. First, we provide a specific example of how hypsometry must be considered in analyses of alpine glacial dates; specifically, higher altitude sources areas will likely lead to a wider range of climate conditions conducive to glacial advance. Our results provide an explanation for some observed asynchrony by demonstrating that neighboring glaciers can have opposite responses to a uniform change in climate. Second, the commonly used technique of dating moraines in a single valley and then correlating that date to other moraines in the area that show a similar ELA depression should be conducted cautiously. Others have noted this [e.g., Benn et al., 2005], but here we provide quantitative evidence that different climate regimes can lead to nearly identical ELAs. Third, our data yield additional insights into Himalayan paleoclimate. The amassed paleoclimate data show the early Holocene was wetter than the pre-Holocene period, but few terrestrial data sources exist, particularly for paleotemperature. The lower ELA observed in the late-glacial, despite less snow, strongly suggests that this period was colder, not just drier than the early Holocene.

Despite the potential of using moraines as paleoclimatic indicators, the complexity of glacial behavior has long challenged researchers. In this study, we demonstrate how hypsometry influences glacial behavior. In addition we present the largest and best-constrained 10Be dates available for glacial advances in the central or western Nepal Himalaya.

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